Statistical and software tools for quantifying uncertainty in event attribution

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Outline

- Statistical methods for event attribution
 - Sources of uncertainty
 - Risk ratio estimation methods
 - Methods for quantifying uncertainty
 - Recommendations
- Illustration with Texas heatwave of 2011
 - Sensitivity analysis with respect to event definition
 - Use of climextremes package for risk ratio estimation
- climextremes package
 - Illustration on station precipitation data
 - Features/capabilities

Sources of uncertainty in model-based event attribution

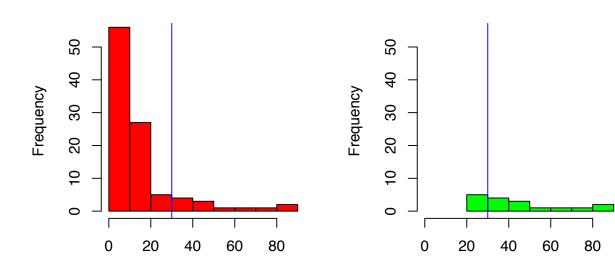
- Sampling uncertainty
 - Due to variability in the earth system
 - Amenable to frequentist or Bayesian statistical treatment
 - Uncertainty decreases with larger ensembles
- Non-sampling uncertainty
 - Sources:
 - Boundary condition uncertainty
 - Model parametric uncertainty
 - Model structural uncertainty
 - Not amenable to frequentist treatment
 - Does not decrease with larger ensembles
 - Possibly characterized based on sensitivity analysis or drawing from prior distribution over boundary conditions, parameter values, models, etc.

Methods to estimate probabilities for risk ratio (RR)

- Fit a parametric statistical model to climate variable values
 - E.g., lognormal or gamma distribution
 - Estimate probability of exceeding the threshold from fitted distribution
 - Strong assumption about appropriateness of model (all data used)
- Count exceedances of threshold (binomial sample) amongst climate variable values
 - "Nonparametric" no distributional assumption
 - More involved to account for dependence (e.g., daily data)
 - High uncertainty when there are very few events, but can be effective for RR
- Fit extreme value distribution (e.g., GEV or peaks-over-threshold (POT)
 - Theoretically justified when threshold is far in tail

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- Informed only by extreme values
- Hard to use with seasonal events because of small sample sizes

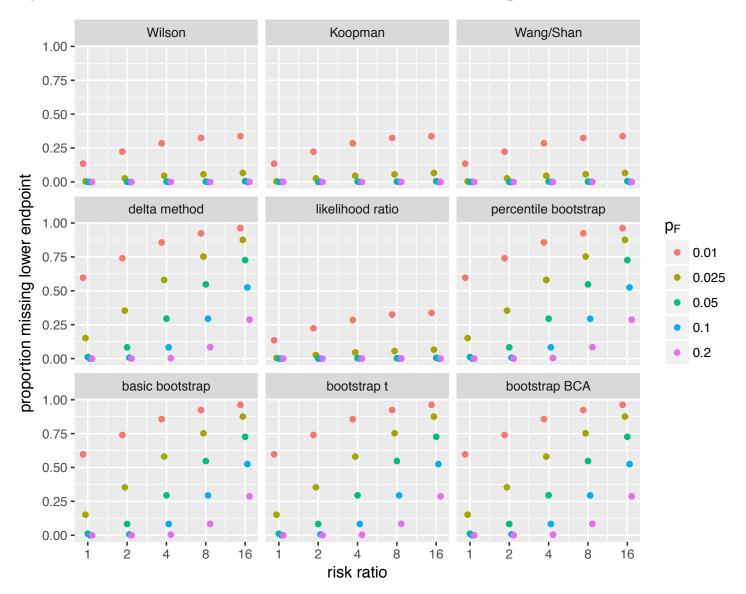


toil values

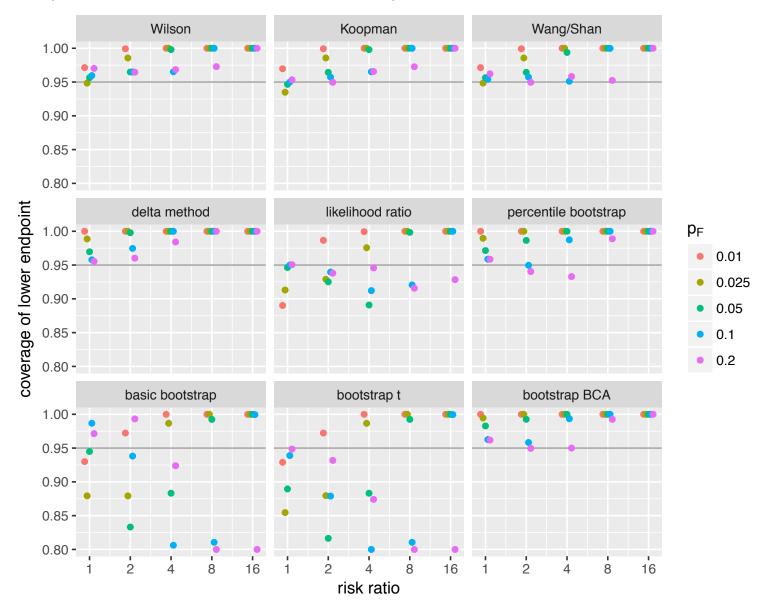
Methods to estimate uncertainty in RR

- Methods for extreme value methods and binomial counting
 - Asymptotic statistical calculations (delta method / propagation of error)
 - Assumption of normality
 - Bootstrap
 - Standard statistical bootstrap gives a confidence interval, not a Bayesian probability interval
 - Estimation of RR in bootstrap samples often fails (lack of EVA convergence, zeros in binomial counting approach)
 - Likelihood ratio-based interval
- Methods for binomial counting (from epidemiology/biostatistics)
 - Wilson's method
 - Koopman's method
 - Wang/Shan method
- All but asymptotic and bootstrap can give interval when RR estimate is 0 or infinity, e.g., (12.8, Infinity)

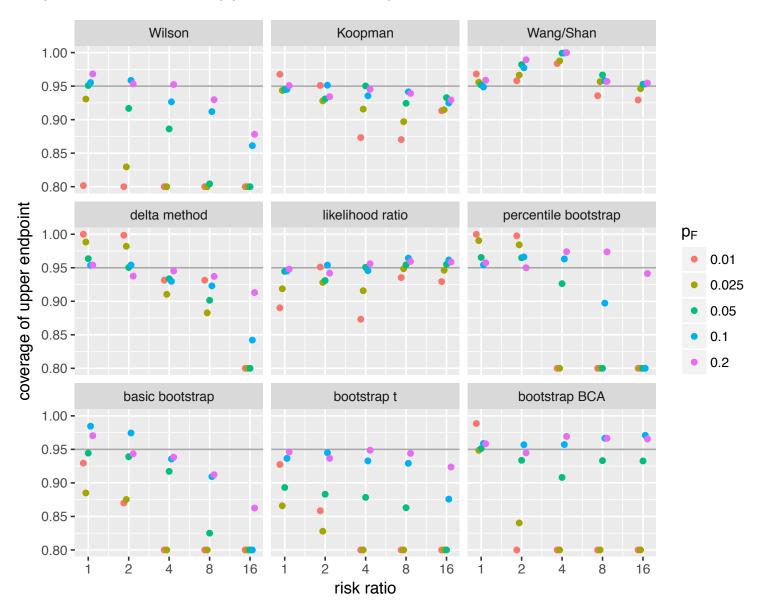
Proportion of times interval cannot be calculated (e.g., 0 events / 0 events)



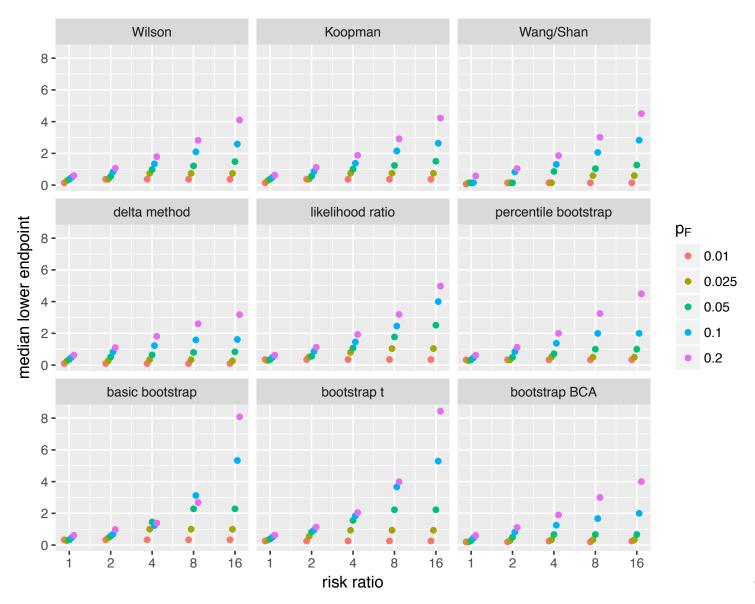
Proportion of times lower interval endpoint includes true RR (95% is best)



Proportion of times upper interval endpoint includes true RR (95% is best)



Average value of lower endpoint (higher is better)

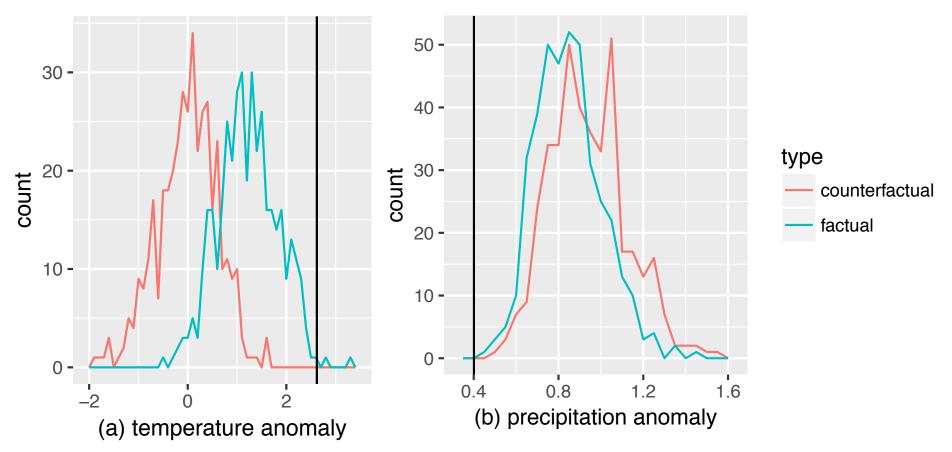


Methods to estimate uncertainty in RR: conclusions

- Bootstrap methods
 - Often fail to provide an interval
 - Poor statistical performance
- Likelihood ratio method
 - Reasonably good statistical performance, but coverage sometimes too low
 - Will work with both extreme value analysis and binomial counting
- Epidemiology/biostatistical methods
 - Work only for binomial counting
 - Koopman and Wang-Shan methods generally perform well
- All methods except Wang-Shan available in climextremes software package

RR analysis example: Texas 2011 heatwave / drought

- CAM5.1 400-member ensembles for factual and counterfactual
- March-August temperature and rainfall over Texas
- Estimation done using climextremes software
- Temperature: 2/0 is RR count-based estimate
- Precipitation: 0/0 is RR count-based estimate



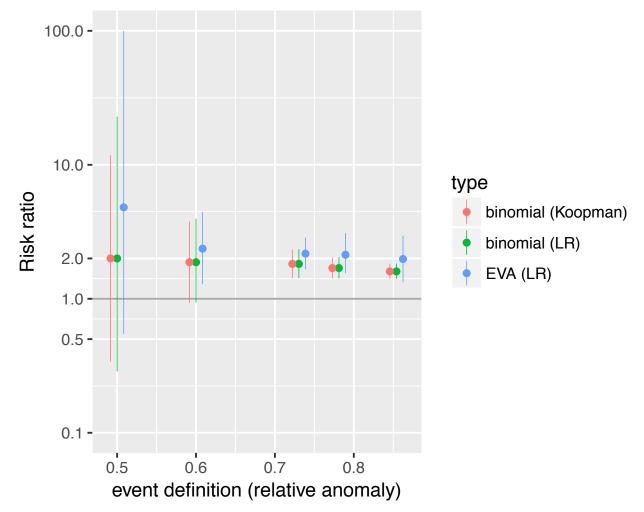
Texas 2011 temperature analysis

Actual event is 2.62 degree anomaly

Event	Number exceedances	RR esti- mate	EVA lik ratio Cl	Binomial lik ratio Cl	Binomial Koopman Cl
2.62 (actual)	2/0	Inf	(12.8, Inf)	(1.0, Inf)	(0.7, Inf)
2.0	43/0	Inf	NA	(31, Inf)	(16, Inf)
1.5	129/3	43	NA	(19, 133)	(17, 108)
1.03 (20-year event)	245/11	22	NA	(14, 38)	(14, 36)
0.73 (10-year event)	314/40	7.9	NA	(6.2, 10.2)	(6.1, 10.1)
0.43 (5-year event)	357/90	4.0	NA	(3.4, 4.7)	(3.4, 4.6)

Texas 2011 precipitation analysis

- Recall that actual event (40% of historical average precipitation) has no events in factual or counterfactual ensembles
- Extreme value analysis (EVA) gives (0.01, Inf) as interval
- EVA not really appropriate for less extreme events but shown anyway

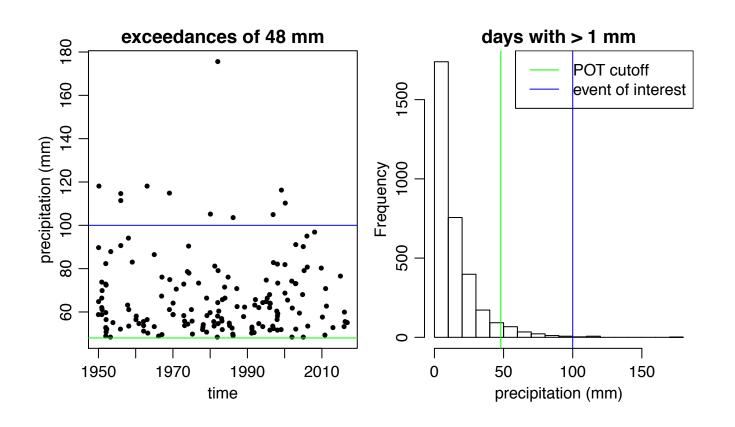


climextremes software

- High-level goals
 - Operate from Python or R
 - Provide risk ratio calculations and extreme value analysis fitting (GEV and POT)
 - Handle common situations with climate data
 - Designed for both observations and model output
- Technical features
 - Use of covariates for any extreme value distribution parameter (nonstationary fitting)
 - Estimation with uncertainty for risk ratios, return values, return periods, differences in return values
 - Various techniques for estimating uncertainty
 - Statistically rigorous estimation with model ensembles
 - Statistically rigorous treatment of missing values (for POT)
 - Allows weighting (e.g., weighting nearby stations)

climextremes example

- US GHCN Santa Cruz precipitation
 - 1950-2016
 - November-May rainy season
 - 270 missing days



climextremes example

Stationary peaks-over-threshold fit

result = climextremes.fit_pot(numpy.array(exc.y), nBlocks = nyears, threshold = threshold, firstBlock = missing.seasonyear[0], blockIndex = numpy.array(exc.seasonyear), index = numpy.array(exc.day), proportionMissing = numpy.array(missing.propMiss), declustering = 'noruns', returnPeriod = 20, returnValue = 100, bootSE = False)

Nonstationary peaks-over-threshold fit

```
# linear location trend in time - # contrast 2015 returnValue and return probability with that for 1950
```

```
resultNS = climextremes.fit_pot(numpy.array(exc.y), x = numpy.array(missing.seasonyear), locationFun = 1, nBlocks = nyears, threshold = threshold, firstBlock = missing.seasonyear[0], blockIndex = numpy.array(exc.seasonyear), index = numpy.array(exc.day), proportionMissing = numpy.array(missing.propMiss), declustering = 'noruns', xNew = 2015, xContrast = 1950, returnPeriod = 20, returnValue = 100, bootSE = False)
```

climextremes example

Stationary peaks-over-threshold fit

```
# 20-year return value and standard error
result['returnValue']
# 120.3 mm
result['se_returnValue'] # return value standard error (asymptotic)
# 7.9 mm
result['logReturnProb'] # log of probability of exceeding 'returnValue=100'
# -1.98
# confidence interval on return probability for 100 mm event
np.exp(result['logReturnProb'] + np.array((-2, 2))*result['se_logReturnProb'])
# (0.0872262, 0.2200104)
```

Nonstationary fit with location linear in year

```
# change in return value (2015 - 1950) and standard error of the change
resultNS['returnValueDiff']
# -2.68 mm
resultNS['se_returnValueDiff']
# 5.37 mm
# risk ratio 2015 / 1950 for 100 mm event with confidence interval
np.exp(resultNS['logReturnProbDiff'])
# 0.88
np.exp(resultNS['logReturnProbDiff'] + np.array((-2, 2))*resultNS['se_logReturnProbDiff'])
# (0.51, 1.49)
```

References / Links

- Statistical methods:
 - Paciorek C.J., D.A. Stone, and M.F. Wehner. 2018. Quantifying statistical uncertainty in the attribution of human influence on severe weather. Weather and Climate Extremes, accepted.
 - https://arxiv.org/abs/1706.03388
- climextremes software (version 0.2.0):
 - Available via conda for python
 - Available via CRAN for R
 - Repository: https://bitbucket.org/lbl-cascade/climextremes-dev